

On the downward continuation of gravitational gradients (GOCE-GDC project)

Novák¹ P., Sebera^{1,2} J., Vařko¹ M., Baur³ O.

1. NTIS, University of West Bohemia, Czech Republic
2. Astronomical Institute, Academy of Sciences of the Czech Republic
3. Space Research Institute, Austrian Academy of Sciences, Austria

GGHS, Venice, Italy

www.gravityfield.org/conference/

9 - 12 October, 2012

1 Introduction

- Insight into downward/upward continuation
- Basic tools

2 Step 1: From real orbit to a mean sphere

- Motivation
- Simulation
- Real-data example

3 Step 2: UDC from mean sphere to zero level

- Overview
- Iterative algorithm
- Noise-free example: Error amplification
- Noise-free example: Edge effect
- Real-data example

4 Concluding remarks

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Insight into upward/downward continuation

Why?

“The main advantage of enhancing potential field by downward continuation with respect to other derivative-based techniques is that the physical dimensions of the transformed field are the same as in the original data” (Fedi and Florio 2002).

But!

“Continuing downward, the functions become rougher. The transition from the flight level to the level of the Earth causes an amplification of the high-frequencies” (Hofmann-Wellenhof and Moritz 1986).

- ⇒ Upward cont. - acts as a low pass filter
- ⇒ Downward cont. - vice versa: usually **strong error amplification**

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Basic options for continuation of potential data

- Gradient method - the function is developed into a Taylor series and continued via $f = f(0) + \frac{f'}{1!}dx + \frac{f''}{2!}dx^2 + \dots$
 - + Very simple, flexible and very fast
 - The derivatives f', f'', \dots must be known
 - Usually applicable for small differentials (dx)
- Poisson integral - standard tool for continuation of potential data
 - + Rigorous (solution of the Dirichlet problem) and quite flexible (curve, plane, sphere, ellipsoid)
 - + Plane - fast, solvable via FFT
 - Curve, Sphere, Ellipsoid - integral equation must be solved
- Other methods - global basis functions (sph. harmonics, ...), local (wavelets, ...), collocation, combined strategies (google for more)

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Step 1: From real orbit to a mean sphere (gradient method)

- Why? It can simplify (speed up) any further GOCE data processing (gridding, SHA, ...)
- Why not? The GOCE orbit \vec{r} is nearly circular: $d\vec{x} = \Delta\vec{r}$ is small ($|\Delta\vec{r}| \leq 16$ km)
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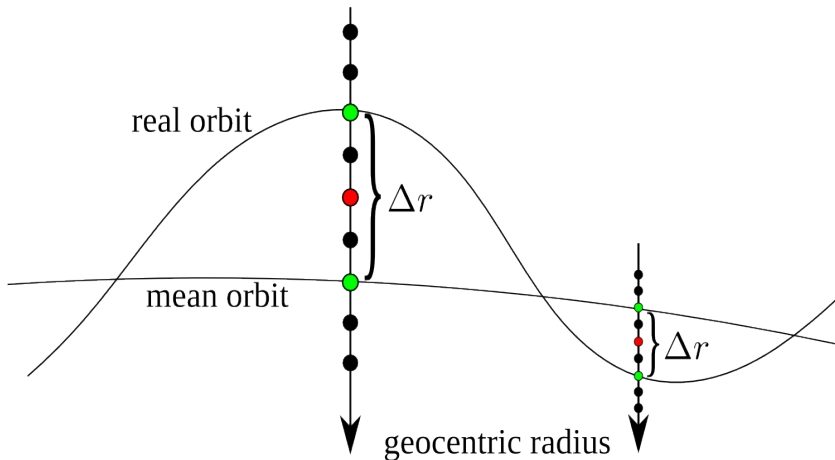
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We have used 9-point stencil for UD continuation of GOCE data

$$V_{ij}^{cont} = V_{ij}^{GOCE} \pm \Delta r \frac{dV_{ij}^0}{dr}$$

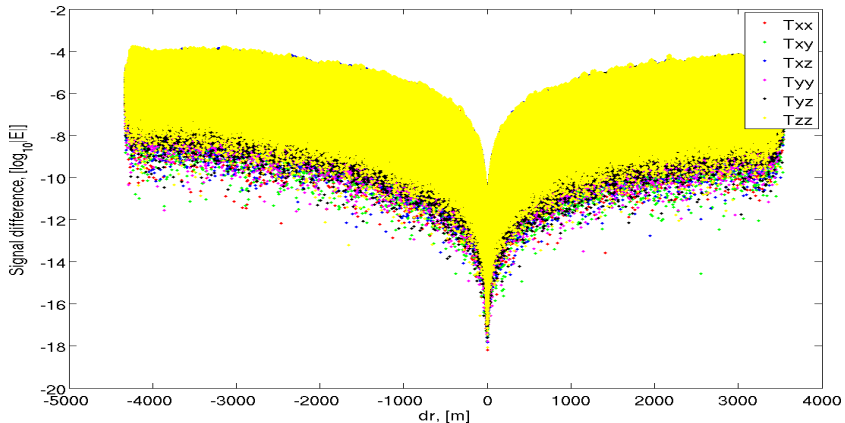
$$\frac{dV_{ij}^0}{dr} = \{aV_{ij}^{-4} - bV_{ij}^{-3} + cV_{ij}^{-2} - dV_{ij}^{-1} + eV_{ij}^0 + dV_{ij}^{+1} - cV_{ij}^{+2} + bV_{ij}^{+3} - aV_{ij}^{+4}\}/dr$$

$$\text{with } a = \frac{1}{280}, b = \frac{4}{105}, c = \frac{1}{5}, d = \frac{4}{5}, e = 0$$

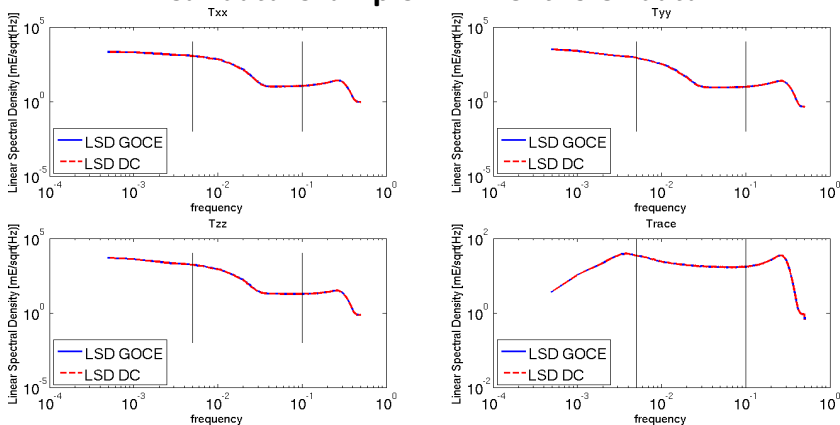


Simulation - height vs. accuracy ($\Delta r = 4dr$)

- Estimated when applied to synthetic data from GOCE-only model
- Minimal accuracy = 0.1 mE with $\text{RMS} \simeq 10^{-6}$ E
- \Rightarrow GOCE data are lucky, they can be UD continued to a mean sphere with gradient method

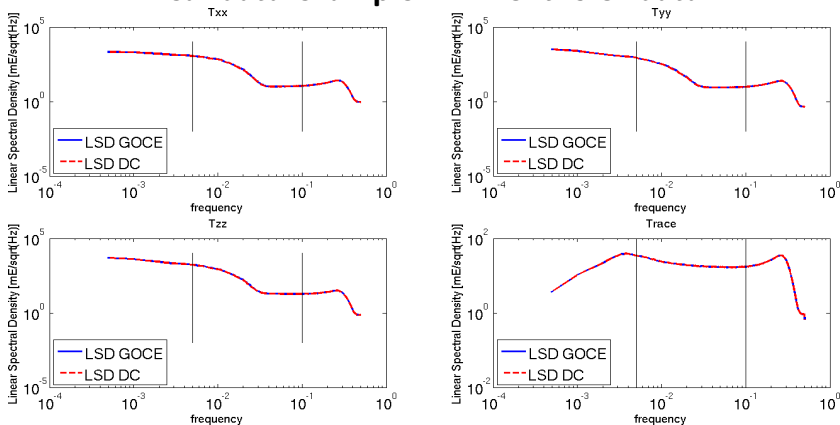


Real-data example - 2 months of data



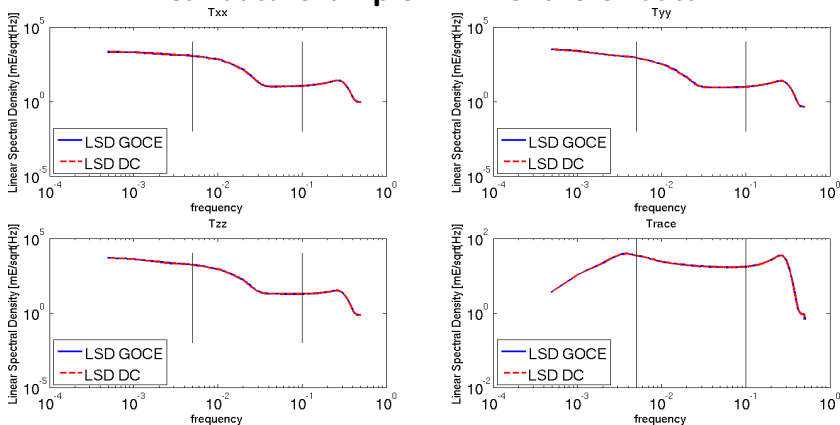
- SDs of the original GOCE (LNOF) and continued data almost the same
- \Rightarrow UDC from real orbit to mean sphere preserves spectral properties
- \Rightarrow Easily applicable to the GRF data

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Step 2: UDC from mean sphere to zero level (Poisson integral)

- Gridding data from the previous step needed (we loose time series!)
- Regular grid is more suitable for surface quadratures
- We have adopted two strategies:
 - ① Iterative approach (Xu, 2007) based on spherical Poisson kernel
 - ② Direct integration with reciprocal spherical Poisson kernel (Novak, 2002)
- Poisson integrals used read $F_k = K_{k,l} F_l$

$$T(P) = \frac{R(r^2 - R^2)}{4\pi} \iint_{\sigma} \frac{T(Q)}{l^3} d\sigma$$

$$T_z(P) = \frac{R(r^2 - R^2)}{4\pi} \frac{R}{r} \iint_{\sigma} \frac{T_z(Q)}{l^3} d\sigma$$

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Iterative algorithm equipped by spherical Poisson integral

- Based on Xu et al. (2007, Geophysical prospecting) but on the sphere

- It applies the upward continuation until the residuals decrease

$$\delta = s_p(F_{\text{input}} - F_i)$$

- It can be applied locally!
- We use the same kernel for all V_{ij} :

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- Before doing so, set:
 - i - number of iterations
 - s_p - "sensitivity" parameter
 - tile/grid size wrt height
 - Algorithm can be split to ladder-like algorithm
 - \Rightarrow many degrees of freedom!

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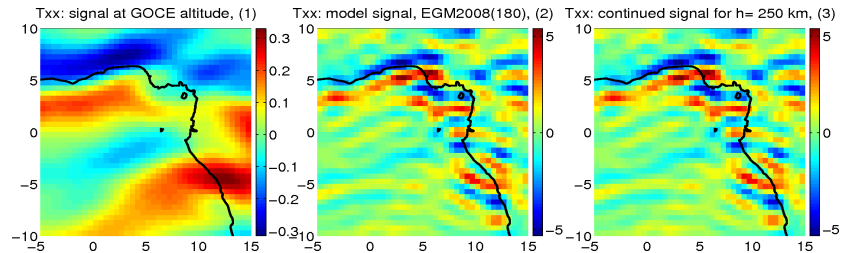
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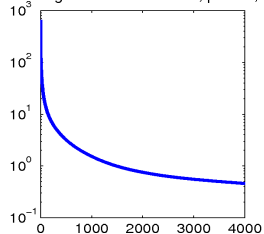
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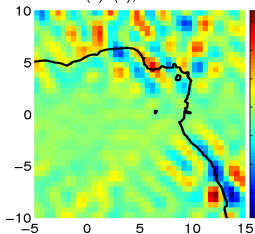
Iterative algorithm: noise-free example



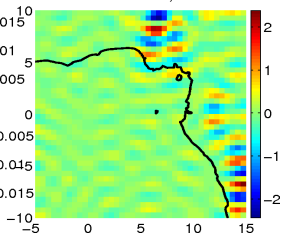
Convergence of the iteration, $ps=1.8$, (4)



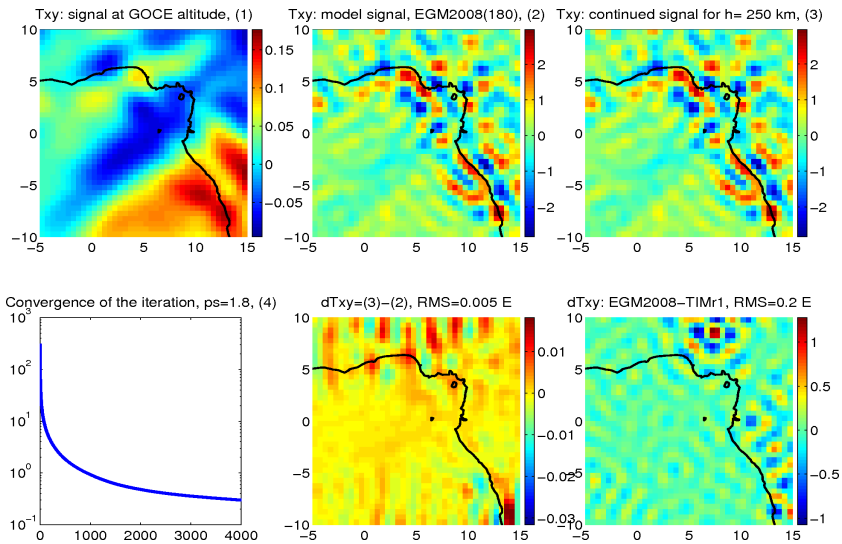
$dTxx=(3)-(2)$, $RMS=0.004$ E



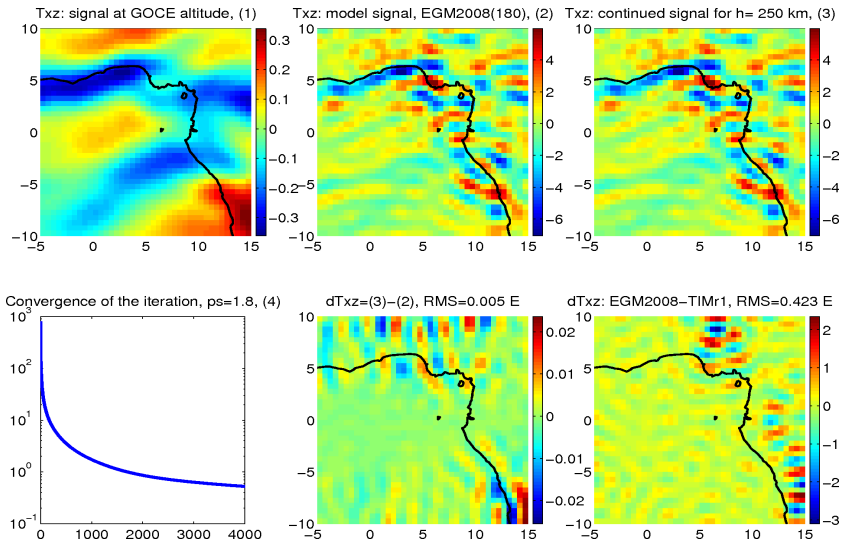
$dTxx: EGM2008-TIMr1$, $RMS=0.374$ E



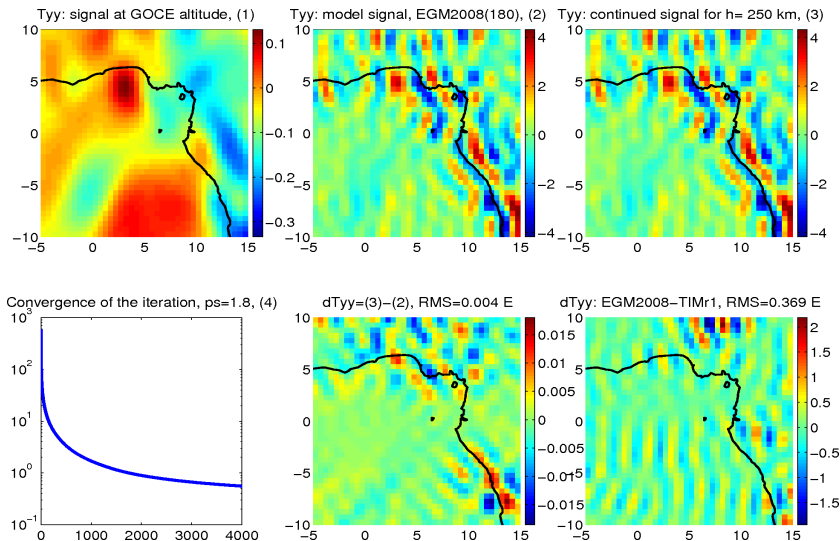
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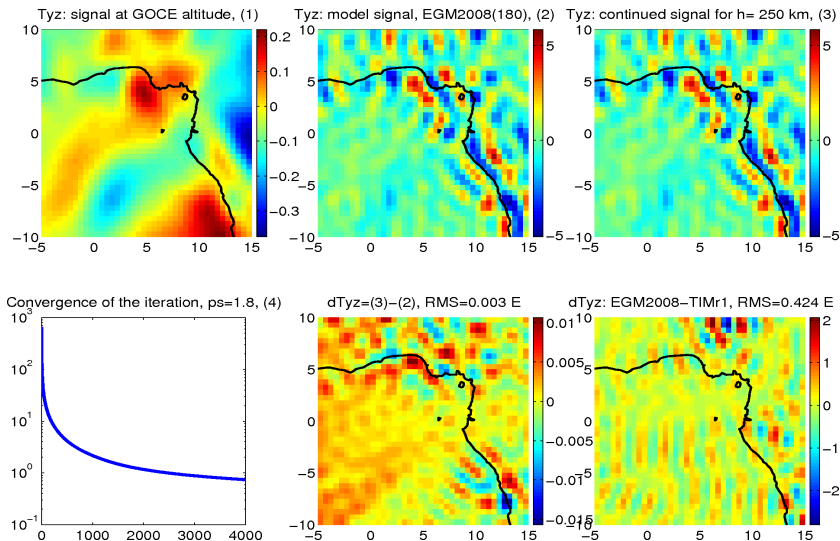
Iterative algorithm: noise-free example



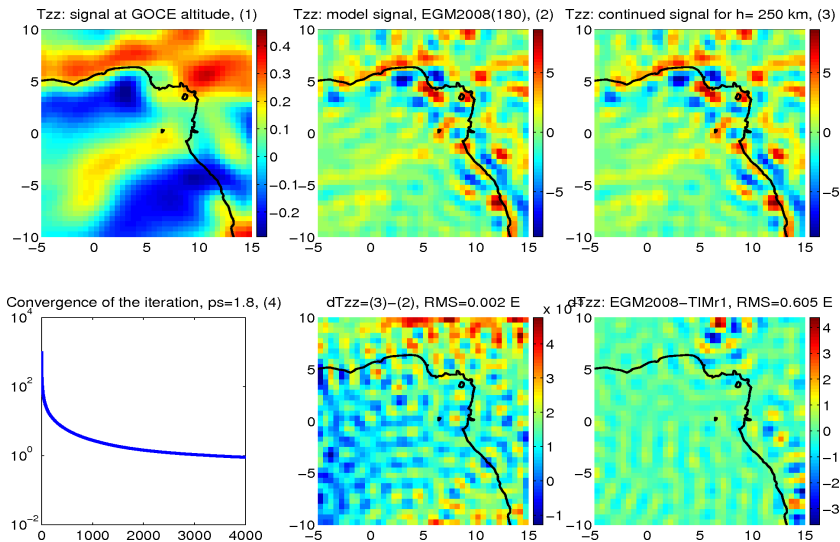
Iterative algorithm: noise-free example



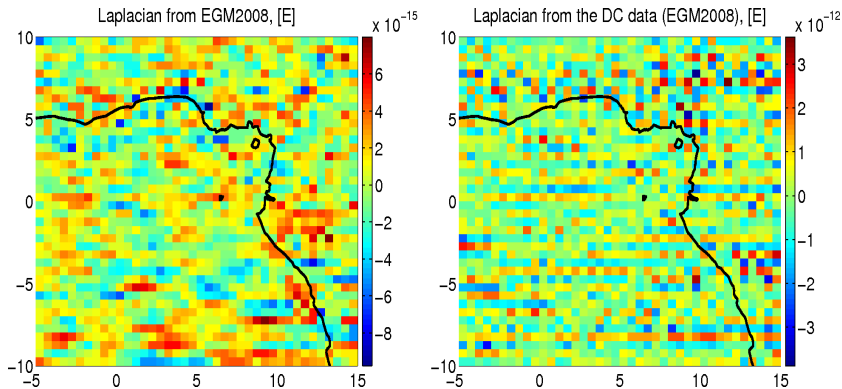
Iterative algorithm: noise-free example



Iterative algorithm: noise-free example

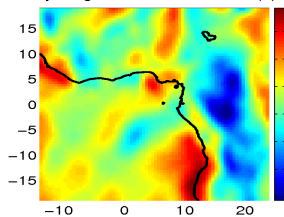


Iterative algorithm: Error amplification (check ΔT)

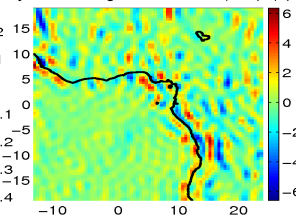


Iterative algorithm: noise-free example (with edge effect)

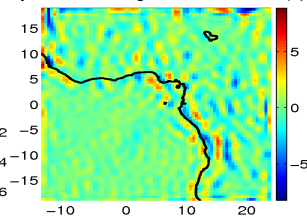
Tyz: signal at GOCE altitude, (1)



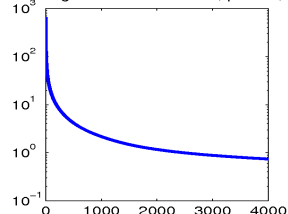
Tyz: model signal, EGM2008(180), (2)



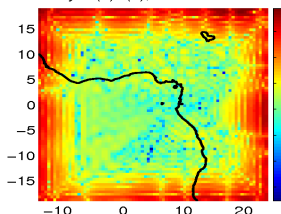
Tyz: continued signal for h= 250 km, (3)



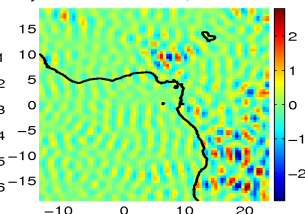
Convergence of the iteration, ps=1.8, (4)



dTyz=(3)-(2), RMS=0.697 E



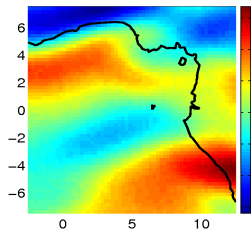
dTyz: EGM2008-TIMr1, RMS=0.492 E



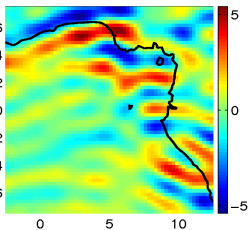
Iterative algorithm: real GOCE data

No filtering

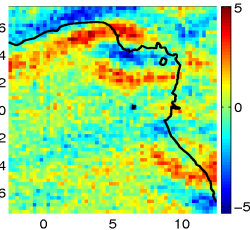
Txx: signal at GOCE altitude, (1)



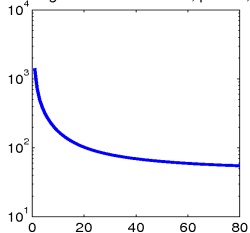
Txx: model signal, EGM2008(180), (2)



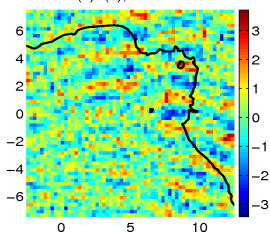
Txx: continued signal for h= 250 km, (3)



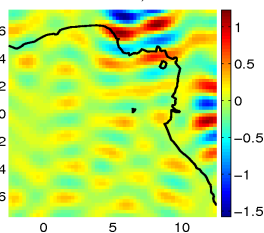
Convergence of the iteration, ps=2.9, (4)



dTxx=(3)-(2), RMS=0.976 E



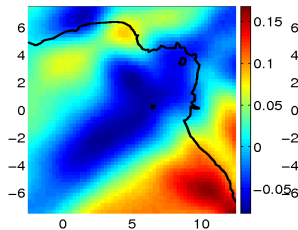
dTxx: EGM2008-TIMr1, RMS=0.282 E



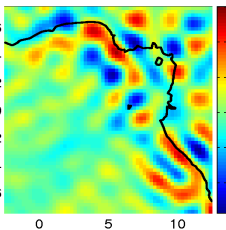
Iterative algorithm: real GOCE data

No filtering

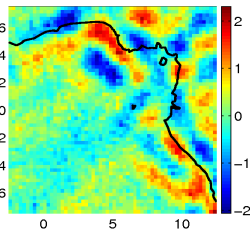
Txy: signal at GOCE altitude, (1)



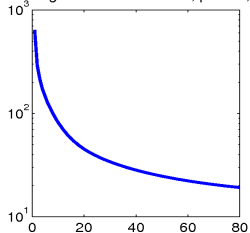
Txy: model signal, EGM2008(180), (2)



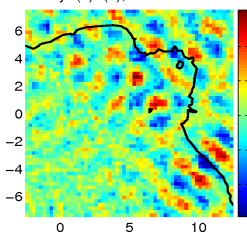
Txy: continued signal for h= 250 km, (3)



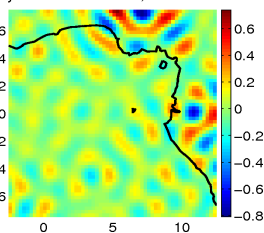
Convergence of the iteration, ps=2.9, (4)



dTxy=(3)-(2), RMS=0.403 E



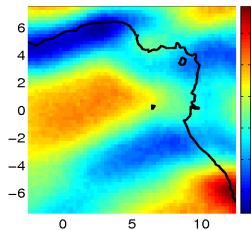
dTxy: EGM2008-TIMr1, RMS=0.154 E



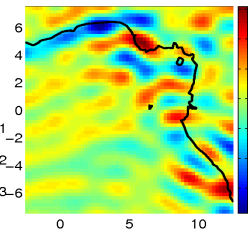
Iterative algorithm: real GOCE data

No filtering

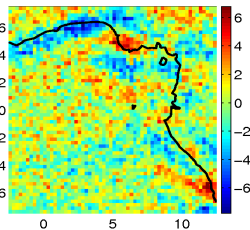
Tzx: signal at GOCE altitude, (1)



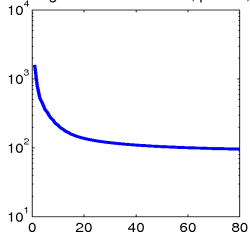
Tzx: model signal, EGM2008(180), (2)



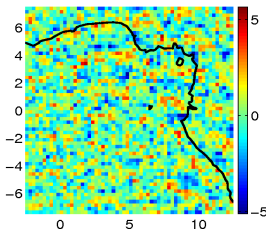
Tzx: continued signal for h= 250 km, (3)



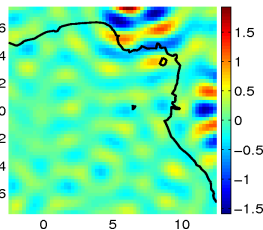
Convergence of the iteration, ps=2.9, (4)



dTzx=(3)-(2), RMS=1.532 E



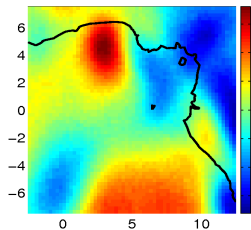
dTzx: EGM2008-TIMr1, RMS=0.322 E



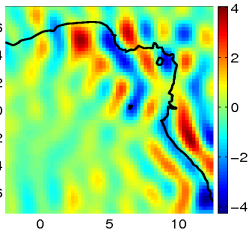
Iterative algorithm: real GOCE data

No filtering

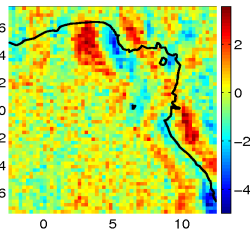
Tyy: signal at GOCE altitude, (1)



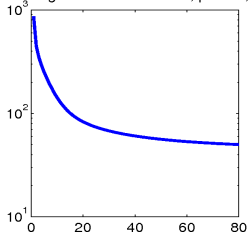
Tyy: model signal, EGM2008(180), (2)



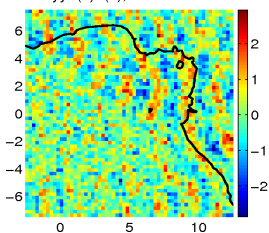
Tyy: continued signal for h= 250 km, (3)



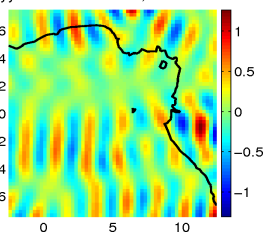
Convergence of the iteration, ps=2.9, (4)



dTyy=(3)-(2), RMS=0.825 E



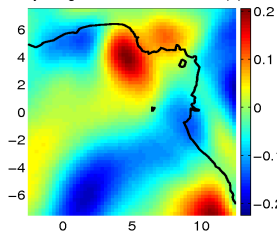
dTyy: EGM2008-TIMr1, RMS=0.316 E



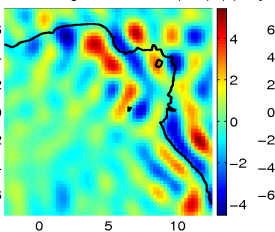
Iterative algorithm: real GOCE data

No filtering

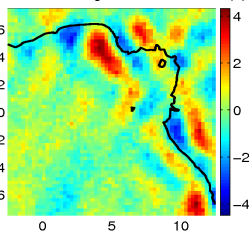
Tyz: signal at GOCE altitude, (1)



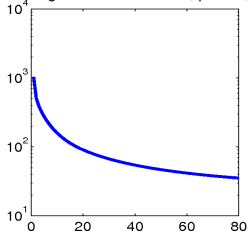
Tyz: model signal, EGM2008(180), (2)



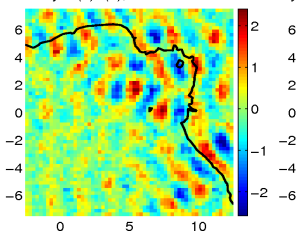
Tyz: continued signal for h= 250 km, (3)



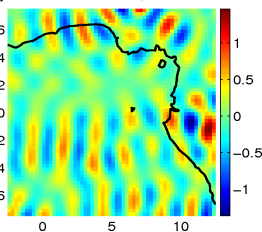
Convergence of the iteration, ps=2.9, (4)



dTyz=(3)-(2), RMS=0.685 E



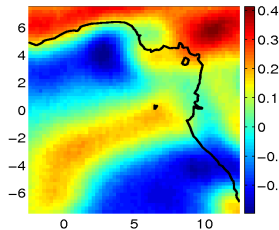
dTyz: EGM2008-TIMr1, RMS=0.347 E



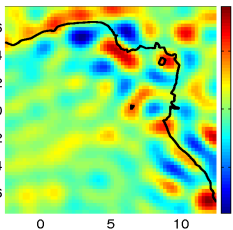
Iterative algorithm: real GOCE data

No filtering

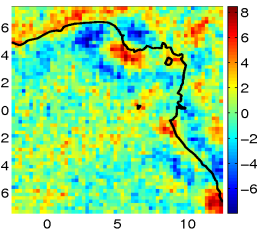
Tzz: signal at GOCE altitude, (1)



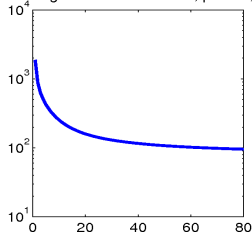
Tzz: model signal, EGM2008(180), (2)



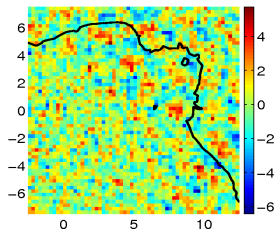
Tzz: continued signal for h= 250 km, (3)



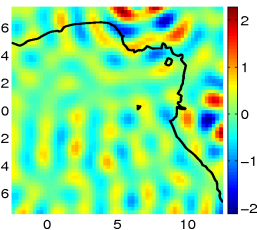
Convergence of the iteration, ps=2.9, (4)



dTzz=(3)-(2), RMS=1.605 E



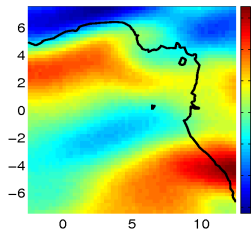
dTzz: EGM2008-TIMr1, RMS=0.477 E



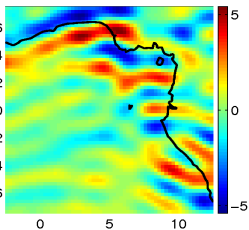
Iterative algorithm: real GOCE data

No filtering

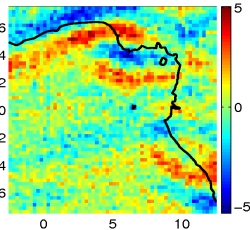
Txx: signal at GOCE altitude, (1)



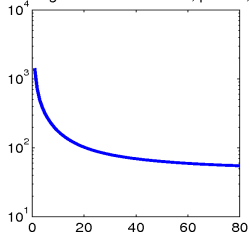
Txx: model signal, EGM2008(180), (2)



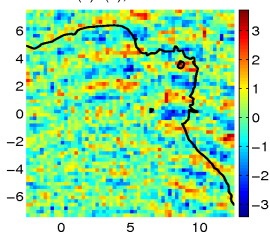
Txx: continued signal for h= 250 km, (3)



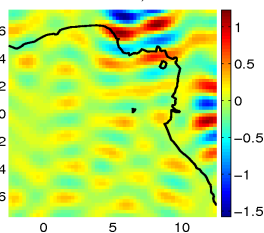
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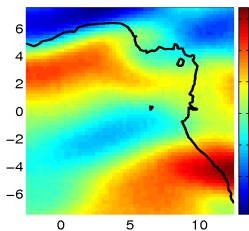
dTxx: EGM2008-TIMr1, RMS=0.282 E



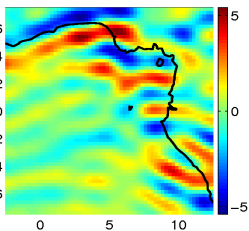
Iterative algorithm: real GOCE data

Filtered (anisotropic filter)

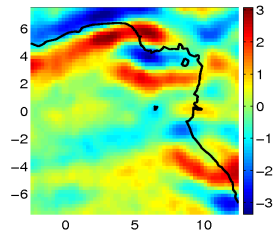
Txx: signal at GOCE altitude, (1)



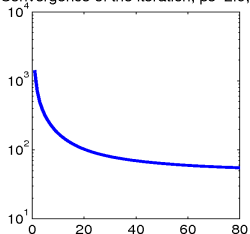
Txx: model signal, EGM2008(180), (2)



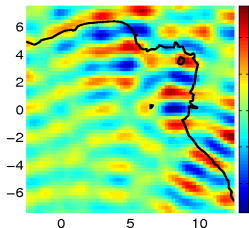
Txx: continued signal for h= 250 km, (3)



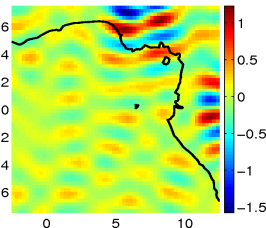
Convergence of the iteration, ps=2.9, (4)



dTxx=(3)-(2), RMS=0.846 E



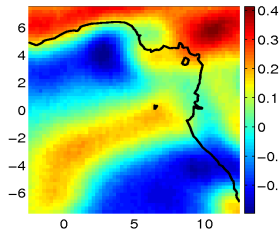
dTxx: EGM2008-TIMr1, RMS=0.282 E



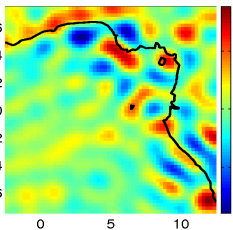
Iterative algorithm: real GOCE data

No filtering

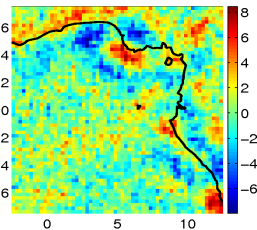
Tzz: signal at GOCE altitude, (1)



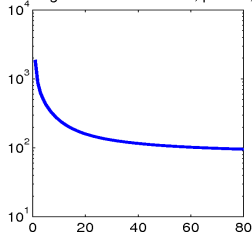
Tzz: model signal, EGM2008(180), (2)



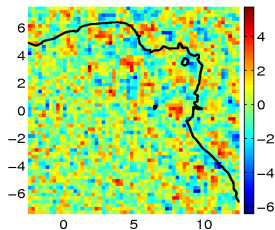
Tzz: continued signal for h= 250 km, (3)



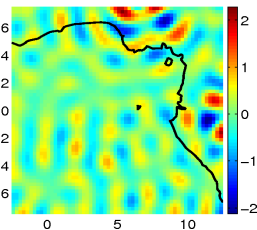
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dTzz=(3)-(2), RMS=1.605 E



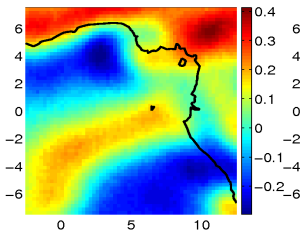
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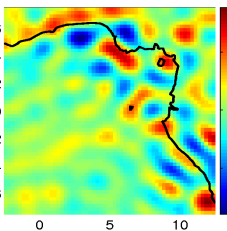
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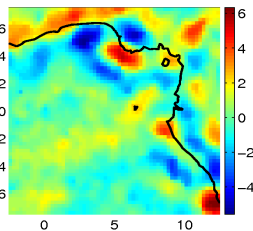
Tzz: signal at GOCE altitude, (1)



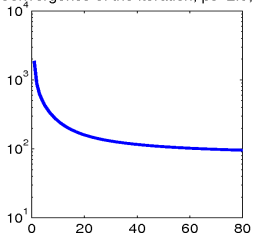
Tzz: model signal, EGM2008(180), (2)



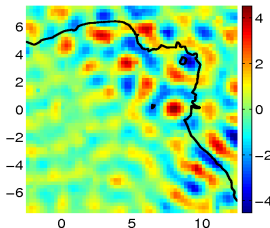
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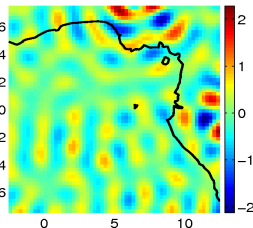
Convergence of the iteration, ps=2.9, (4)



dTzz=(3)-(2), RMS=1.208 E



dTzz: EGM2008-TIMr1, RMS=0.477 E



Concluding remarks on downward continuation of GOCE data (1. From real to mean sphere)

- GOCE orbit: $\delta r \in \langle -16, 16 \rangle$ km
- Simple gradient method can be applied to data in any frame (GRF, LNOF)
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Concluding remarks on downward continuation of GOCE data (2. From mean sphere to ground level)

- + DC of the GGs for $h = 250$ km is possible! Basic features recovered, for $N = 180$ we got an agreement at 0.5 - 1 E level (best RMS achieved for $N = 130 - 150$)
- + Algorithm performs even at mE level for noise-free data (tested with EGM2008)
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- ? Optimal DC empirical parameters must be found for a given tile and spacing
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